Hydropedological behavior of a chromic vertisol under different plant covers

Comportamiento hidropedológico de un vertisol crómico bajo diferentes coberturas vegetales
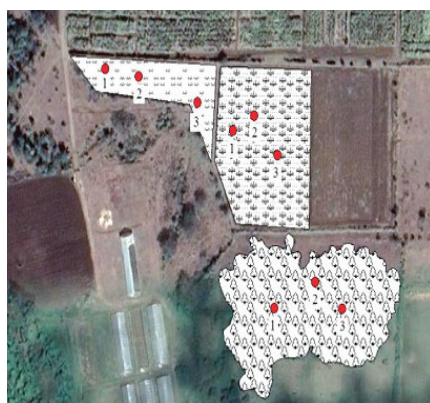
Juan-Alejandro Villazón-Gómez[1]

ABSTRACT

Vertisols present edaphic limiting factors primarily of a physical nature. The study of their hydropedology, supported by multivariate techniques, is of vital importance. The work was developed with the objective of determining, through multivariate analysis, the hydropedological behavior of a Chromic Vertisol under different plant covers. In areas under natural grass, sugarcane, and secondary forest, belonging to the Holguín Sugarcane Provincial Research Station, three random points were chosen in each. Soil penetration resistance, soil bulk density, gravimetric moisture, and volumetric moisture at depths of 0-10, 10-20 and 20-30cm were determined. A Principal Component Analysis and Canonical Correlations were performed using Statistica 7 and Statgraphics Plus XV.II. The first two components made the greatest contribution to the variance, with 83.09%. The greatest contributions (1st component) were given by moisture at all depths and by soil compaction at a depth of 20-30cm. The 2nd component was influenced by soil compaction in the 0-10 and 10-20cm layers. There was a contrast between the variables that characterize the solid phase with which they describe the liquid phase of the soil. There was a correlation between soil moisture and compaction. The first two pairs of canonical variables showed a strong linear correlation and regularly dispersed along the central values of the model, with a remarkable grouping by depths.

Keywords: Principal component analysis; canonical correlations; soil bulk density; soil compaction; soil moisture; soil penetration resistance.

RESUMEN

Los Vertisoles presentan factores edáficos limitantes fundamentalmente de carácter físico. El estudio de su hidropedología, apoyado en técnicas multivariadas, es de vital importancia. El trabajo se desarrolló con el objetivo de determinar, mediante análisis multivariado, el comportamiento hidropedológico de un Vertisol Crómico bajo diferentes usos de la tierra. En áreas bajo pasto natural, caña de azúcar y bosque secundario,

pertenecientes a la Estación Provincial de Investigaciones de la Caña de Azúcar de Holguín, se escogieron tres puntos al azar en cada una. Se determinó la resistencia a la penetración, la densidad aparente, la humedad gravimétrica y la humedad volumétrica a las profundidades de 0-10, 10-20 y 20-30 cm. Se realizó un Análisis de Componentes Principales y de Correlaciones Canónicas, con los softwares Statistica 7 y Statgraphics Plus XV.II. Se encontró que las dos primeras componentes realizaron la mayor contribución a la varianza, con un 83.09%. Los mayores aportes (1ra componente) estuvieron dados por la humedad en todas las profundidades y por la compactación del suelo en la profundidad de 20-30 cm. La 2da componente estuvo influenciada por la compactación del suelo en las capas de 0-10 y de 10-20 cm. Existió contraposición entre las variables que caracterizan la fase sólida con las que describen la fase líquida del suelo. Existió correlación entre la humedad del suelo y la compactación. Los dos primeros pares de variables canónicas mostraron una fuerte correlación lineal y dispersada regularmente a lo largo de los valores centrales del modelo, con un notable agrupamiento por profundidades.

Palabras clave: Análisis de Componentes Principales; Correlaciones canónicas; Densidad aparente; compactación de suelos; Resistencia a la penetración; humedad del suelo.

INTRODUCTION

Soil is a system composed of three phases: solid, liquid and gas. The arrangement of the solid phase or soil matrix determines the geometry of the pore space in which water and air function (Cid-Lazo *et al.*, 2004b). Therefore, it is important, when it comes to knowing the biophysical environment in which the root development of plants occurs, to carry out the physical characterization of the soil (Batey, 2009). This is a very complex process comparable to those studies aimed at grouping properties that delimit a certain behavior of soil functioning (Cid-Lazo *et al.*, 2004b).

With the change in land use, the first modification of a natural ecosystem is compaction (Torres-Guerrero *et al.*, 2016). Soil compaction causes an increase in mechanical density, destroys and weakens its structure, thereby reducing the porosity of soil aeration. Its effects are reflected in a lower development of the root system of plants which implies a lower vegetative development and impairment of agricultural yields (Gutiérrez-Rodriguez *et al.*, 2014).

Vertisols occupy an area of 335 million ha worldwide, generally in semiarid tropical areas, with annual rainfall ranging between 500-1000 mm. Large areas remain in disuse or are only dedicated to extensive uses (livestock, woodcutting, charcoal burning or similar work) that range from smallholdings that produce crops after the rainy season to small and large-scale irrigated agriculture (IUSS Working Group WRB, 2015).

In Cuba they cover an area of about 695,000 ha (Hernández-Jiménez *et al.*, 2014). They are intensively dedicated to the cultivation of sugar cane and rice. Livestock production occupies 50% of agricultural areas (Cid-Lazo *et al.*, 2016).

These soils are characterized by presenting a layer 25 cm or thicker in the first 100 cm of depth, with slickenside or wedge-shaped aggregates inclined longitudinally between 10-60 ° with respect to the horizontal. 30% or more of clays are in the fine earth fraction in the 0-18 cm deep layer or in the Ap horizon and cracks that open and close periodically (Soil Survey Staff, 2010).

They have the peculiarity that, when they yield or absorb water, their apparent volume changes, with the consequent reorganization of the material (Cid-Lazo *et al.*, 2004a). In this way, the apparent density of the soil

constitutes an indicator of how the physical properties affect the dynamics of water in the soil, its availability and use by plants (Yang *et al.*, 2016).

In this way, the concept of physical fertility is inseparable from the general fertility of the soil (Cid-Lazo *et al.*, 2006). Likewise, when evaluating changes in the physical quality of soils, the use of multivariate statistics allows a better perception and interpretation of the interactions between the variables involved (Mota *et al.*, 2014). Furthermore, in recent years, hydropedology has become a multidisciplinary science that integrates the fields of research corresponding to pedology and hydrology (Lin *et al.*, 2006).

For the aforementioned reasons, the objective of this work is to determine, through multivariate analysis, the hydropedological behavior of a Chromic Vertisol under different plant covers.

MATERIALS AND METHODS

The work was developed, in 2019, in areas of the Guaro Experimental Block (centroid: 20°40′8.46″ of latitude N and 75°46′9.58″ of longitude W), belonging to the Provincial Research Station of the Sugarcane (EPICA) from Holguín, Cuba, at 17.4 m a.s.l. The average annual rainfall is 1,067.6 mm, the monthly average temperature is 25.6°C. This region has a tropical rainy climate; it has dry winters and humid summers; registered as Aw, according to the Köppen climate classification. (Peel *et al.*, 2007).

Three types of plant covers were analyzed (natural pasture, sugarcane and secondary forest) established on a Calcium Gleyic Chromic Vertisol (Hernández-Jiménez *et al.*, 2015). The areas of the three vegetation covers were digitized using the MAPINFO 12.0 software in independent layers. Three random points within the limits of each one were generated using the QGIS 3.10 software (Figure 1). In these points, the soil penetration resistance, soil bulk density, gravimetric moisture, and volumetric moisture were determined.

To determine the soil penetration resistance, an impact penetrometer model IAA / Planalsucar-Stolf was used, with the impacting mass of four kg set at a height of 0.40 m. The transformation of the number of impacts per dm to megapascals (MPa) was carried out through the mathematical expression proposed by Stolf (1991):

$$RP(MPa) = 0.547 + 0.675 * N \qquad (1)$$

Where *RP* is soil penetration resistance, and *N* is the number of impacts per dm. These results were expressed in constant intervals (0.10 cm), up to 30 cm deep.

Undisturbed samples were taken at the same depths mentioned above with cylinders of 105.35 cm^3 in volume. The wet soil mass was determined, and the samples were placed in a stove at 105°C until they reached a constant weight. With the data obtained, soil bulk density was determined (Cid-Lazo *et al.*, 2004a; Cid-Lazo *et al.*, 2006):

$$D_a = \frac{mss}{V}$$

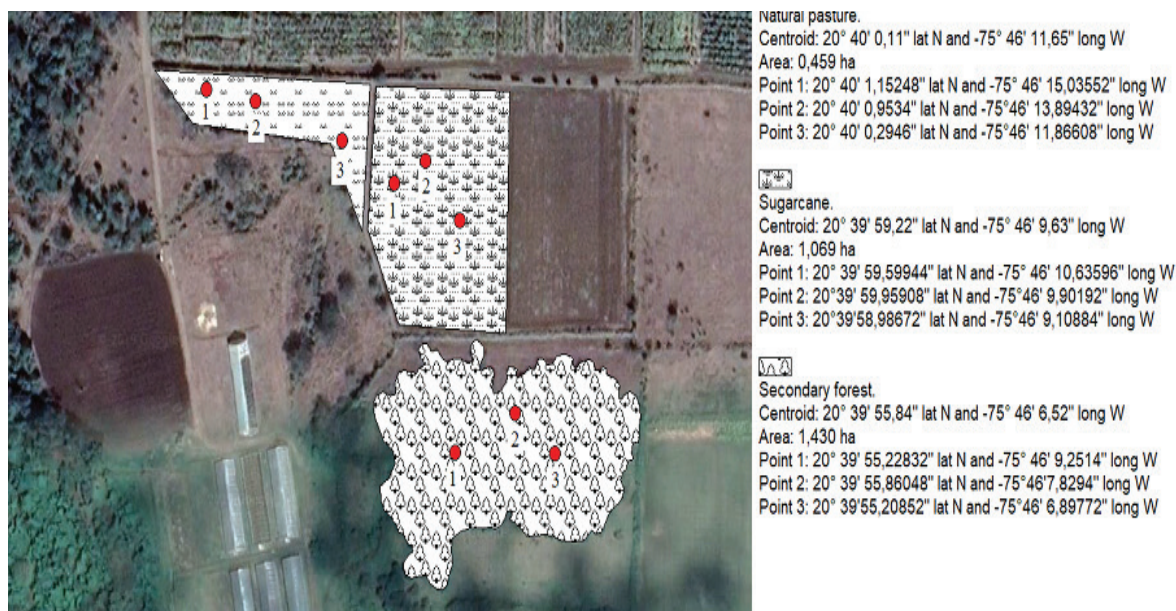where *Da* is soil bulk density, *mss* is dry soil mass, and *V* is the cylinder volume.

Figure 1. Sampling carried out in the three plant covers.

Gravimetric moisture was also obtained, using the equation proposed by Cid-Lazo *et al.* (2004a) and Cid-Lazo *et al.* (2006):

$$\theta_g = \frac{msh - mss}{mss}$$

where θg is the gravimetric moisture, *msh* is the wet soil mass, and *mss* is the dry soil mass. The volumetric moisture was determined from equations (1) and (2) (Cid-Lazo *et al.*, 2004a; Cid-Lazo *et al.*, 2006):

A principal component analysis (PCA) was performed on the non-partitioned data matrix to identify the variables that explain the variance of the data. Subsequently, an analysis of canonical correlations was carried out to verify associations between two qualitative sets of variables (set_1: gravimetric moisture and volumetric moisture, which represent soil moisture; set_2: soil bulk density and soil penetration resistance, which define soil compaction). Statistica 7 and Statgraphics Centurión XV.II software were used for statistical analysis.

RESULTS AND DISCUSSION

The 1st and 2nd components explain 71.97 and 11.12% of the variance, respectively, for a cumulative percentage of 83.09%, as a result of the linear combination of the 12 variables studied. The eigenvalues of the two components were above 1 (See Table 1).

The greatest variation (1st component) given, fundamentally, by the properties related to moisture in the first 20 cm and by hydropedology in the depth of 20-30 cm. As for the 2nd component groups, variation is given by the properties related to soil compaction in the first 20 cm of depth. The 1st component is 6.47 times greater than the 2nd. In both components, there is a contrast between the hydropedological properties that represent soil moisture and those that determine soil compaction.

Table 1. Eigenvectors of the principal component analysis.

Variables	1st component	2nd component
Gravimetric moisture 0-10 cm	**0.832365**	0.437736
Soil bulk density 0-10 cm	-0.130912	**-0.957134**
Volumetric moisture 0-10 cm	**0.897024**	0.029898
Soil penetration resistance 0-10 cm	-0.385339	**-0.801590**
Gravimetric moisture 10-20 cm	0.698635	0.568651
Soil bulk density 10-20 cm	-0.386745	**-0.871434**
Volumetric moisture 10-20 cm	**0.721587**	0.366494
Soil penetration resistance 10-20 cm	-0.586691	-0.772924
Gravimetric moisture 20-30 cm	**0.841134**	0.449639
Soil bulk density 20-30 cm	**-0.744081**	-0.314113
Volumetric moisture 20-30 cm	**0.797638**	0.419031
Soil penetration resistance 20-30 cm	**-0.833466**	-0.412994
Eingenvalue	8.635828	1.334533
Total percentage of variance (%)	71.97	11.12
Cumulative eingenvalue	8.635828	9.970361
Cumulative percentage of variance (%)	71.97	83.09

In the case of the 1st component, as its correlation with moisture is positive, increasing the value of this component increases the water content in the soil, which does not occur with soil compaction in the 20-30cm layer (negative correlation). With respect to the 2nd component, as its value increases, a decrease in soil compaction will occur in the first 20cm of depth (negative correlation).

Figure 2 shows that there are two large groups, one formed by soil moisture and the other by soil compaction. Furthermore, a direct (positive) relationship is observed between soil bulk density and soil penetration resistance. On the contrary, an inverse (negative) relationship was found between soil compaction (represented by soil bulk density and soil penetration resistance) and soil moisture (gravimetric and volumetric moisture). This contrast is more notable

between said water content in the soil and the soil bulk density and soil penetration resistance in a depth of 20-30cm.

The changes caused by soil moisture in soil bulk density are given by the high content of montmorillonitic clays of Vertisols, which causes these soils to contract and expand constantly under conditions of alternating humidity. This fact affects the variations of soil bulk density.

The evidence suggests that the increase in soil compaction, with the consequent reduction of the pore space, hinders the entry of water into the soil, its movement through it, and its storage in a form available to plants. This reduction in the volume of water that can be used by the plants is accompanied by a decrease in root development, which reduces the space explored by the roots in search of water and nutrients present in the soil.
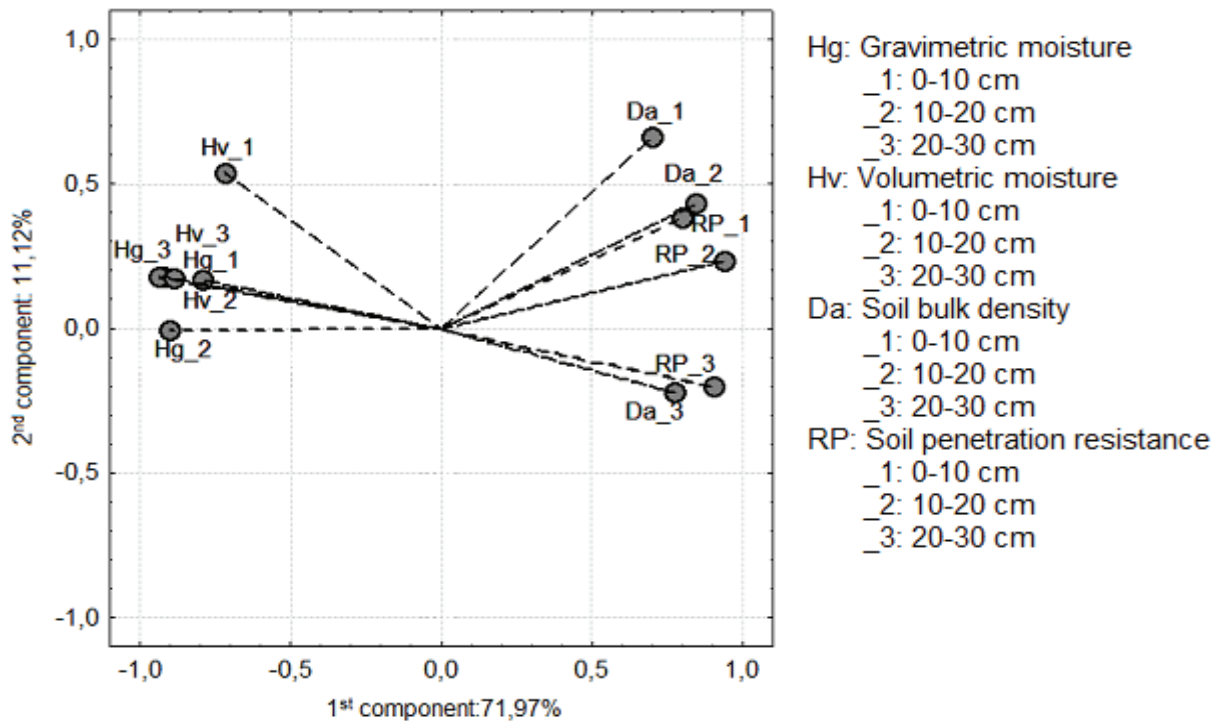
Figure 2. Physical properties dispersion in the principal component analysis. *Da*: soil bulk density; *RP*: soil penetration resistance; *Hg*: gravimetric moisture; *Hv*: volumetric moisture.

Hamza and Anderson (2005) state that the increase in moisture favors the decrease in the volumetric density of the soil, since the water molecules in the pore space of the soil prevent the approach of the soil particles. As the moisture decreases, the soil contracts, thus decreasing its volume and increasing its density. Clay soils turn out to be very susceptible to these wetting and drying cycles.

Gao *et al*. (2016) found, in the first 50 cm of depth, an increase in soil bulk density and the volume of micropores from the surface towards the lower layers of the profile of a soil classified as Vertisol, in Lishu county, Jilin province (China). The increase in soil bulk density caused, especially in non-tillage areas with low soil moisture content, an increase in soil penetration resistance.

In the same way, Jorbenadze *et al*. (2017), when studying the physical properties of the soils of Georgia, found that Vertisols showed an increase in soil bulk density with depth. This caused a decrease in the maximum capacity of the soil to store water when all pore volumes are full of water (saturation water); of the amount of water contained in the soil previously saturated after the excess has been drained by the gravity force action, and the percolation speed has decreased considerably (water at field capacity). Decrease too, the moisture content of the soil with which the plants cannot replace enough water to regain their turgor (permanent wilting point), and of the amount of water available for the development of the crops that is between the field capacity and the permanent wilting point (usable water).

The magnitudes of the Canonical Correlation analysis, in which the two canonical functions have highly significant values is depicted in Table 2. In the case of the first function, it is especially strong due to its high canonical correlation (0.980682). Wilk's lambda (λ), close to 0, indicates that the independent variables account for almost all of the variance.

For the second function, the value of its canonical correlation (0.848368) is also high, although Wilk's lambda shows that there is a lower variance accounting for the independent variables. Being 1-λ equivalent to the regression coefficient (Badii *et al.*, 2007), it is considered that for the variables corresponding to the first function $r^2 = 99\%$ and for the second, $r^2 = 72\%$.

From the construction of linear combinations between the variables of each set, where the variables were standardized from the subtraction of the mean and the division between the standard deviation, the largest correlations were found. The first associated canonical function is:

$$U_1 = -3.20661 * gravimetric\ moisture + 2.68268 * volumetric\ moisture \quad (5)$$

$$V_1 = 0.05807 * soil\ bulk\ density - 0.0742318 * soil\ penetration\ resistance \quad (6)$$

The second associated canonical function is:

$$U_2 = -1.35281 * gravimetric\ moisture + 2.21713 * volumetric\ moisture \quad (7)$$

$$V_2 = 1.26806 * soil\ bulk\ density - 1.64984 * soil\ penetration\ resistance \quad (8)$$

It is seen that in both canonical functions, there is a primary relationship between volumetric moisture and apparent density, with some contributions from gravimetric moisture and resistance to penetration.

Figure 3 shows the correlation between soil moisture (set of explanatory variables CVARA_1) and soil compaction (set of explained variables CVARB_1). It can be seen that the relationship is linear and with a regular dispersion around the central values of the model, although with a certain discontinuity for the values found in the depth of 20-30 cm.

In this way, any variation that occurs in the soil moisture content will cause changes in its compaction. Furthermore, the grouping of the correlations between the hydropedological variables by depths and not by plant covers is notable. Therefore, changes in soil moisture and soil compaction are more influenced by the depth of the soil than by plant covers.

Table 2. Magnitude of the relationships that exist between the two sets.

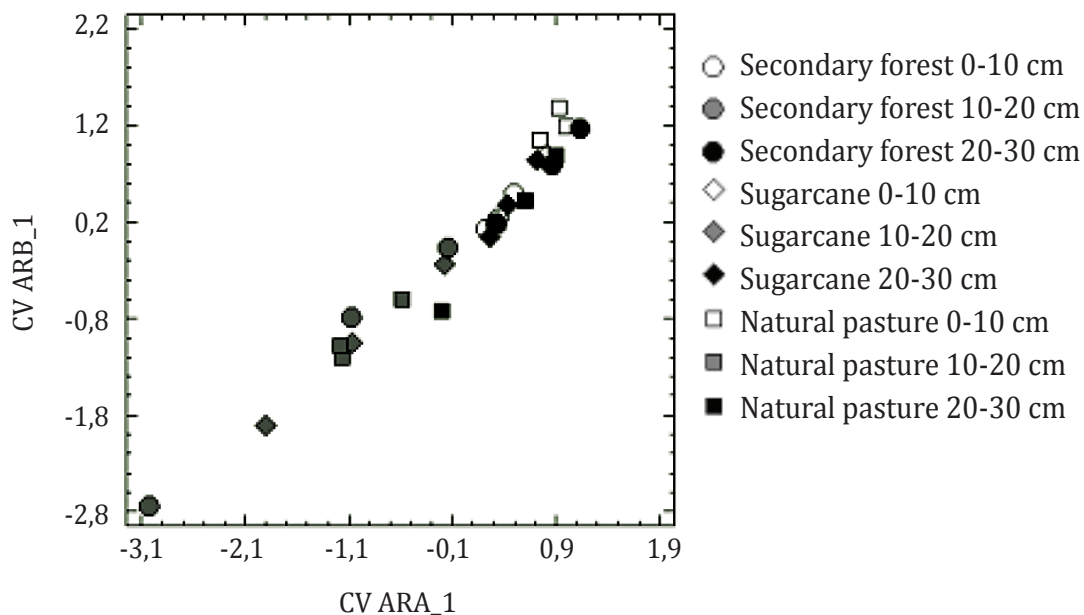| Eingenvalue	Canonical correlation	Wilk' lambda	Chi-square	Deegrees of freedom	P-value
Canonical function 1					
0.961737	0.980682	0.0107241	106.579	4	0.0000
Canonical function 2					
0.719728	0.848368	0.280272	29.892	1	0.0000

Figure 3. Association between soil moisture (CVARA_1) and soil compaction (CVARB_1).

In addition, there are neighborhoods between depths of 0-10 and 20-30cm. The 10-20cm layer is more spaced and with greater discontinuity. This behavior may be due to the fact that the limits between horizon A and B and the highest values of soil compaction are found at this depth.

Krüger *et al*. (2018) point out that there is a correlation between soil penetration resistance and soil moisture content. The latter is a variation factor when the former is measured in soils with different degrees of wetting. In this way, it is useful to determine the soil penetration resistance in analogy with the detection of problems concerning the uptake and movement of water in the soil.

Millán *et al*. (2013), evaluated the effect of the incorporation of biomass on the compaction curve and studied possible alterations in the density and moisture of the soil in a Typical Haplustert located in the municipality of Sincelejo, department of Sucre (Colombia).

They found that, for above the maximum soil bulk density (determined by the Proctor test), the increase in soil moisture caused a decrease in soil compaction. Therefore, the variations in soil bulk density depended on the values of said moisture.

CONCLUSIONS

In Principal Components Analysis, the first two components make the greatest contribution to the variance, with 83.09% of the total variance. The highest contributions (1st component) to the total variance are given by moisture at all depths and by soil bulk compaction at a depth of 20-30 cm. The 2nd component is influenced by the soil compaction in the layers 0-10 and 10-20cm deep. There is a contrast, at all depths, between the variables that characterize the state of the solid phase (soil bulk density and soil penetration resistance) with those that describe the liquid phase of the soil (gravimetric and volumetric moisture).

The Canonical Correlation analysis shows the existence of a correlation between the sets CVARA_1 (gravimetric and volumetric moisture) and CVARB_1 (soil bulk density and penetration resistance). The first two sets of canonical variables show a strong correlation (0.980682 and 0.848368, in each case). This is linear and is regularly dispersed along the central values of the model, with a notable grouping by depth.

There is a strong interrelation between the liquid and solid phases of Vertisols. Variations in soil bulk density and penetration resistance affect the reduction or increase of the moisture content in the pore space of these soils.

Conflict of interest: The authors declare that there is no conflict of interest.

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